

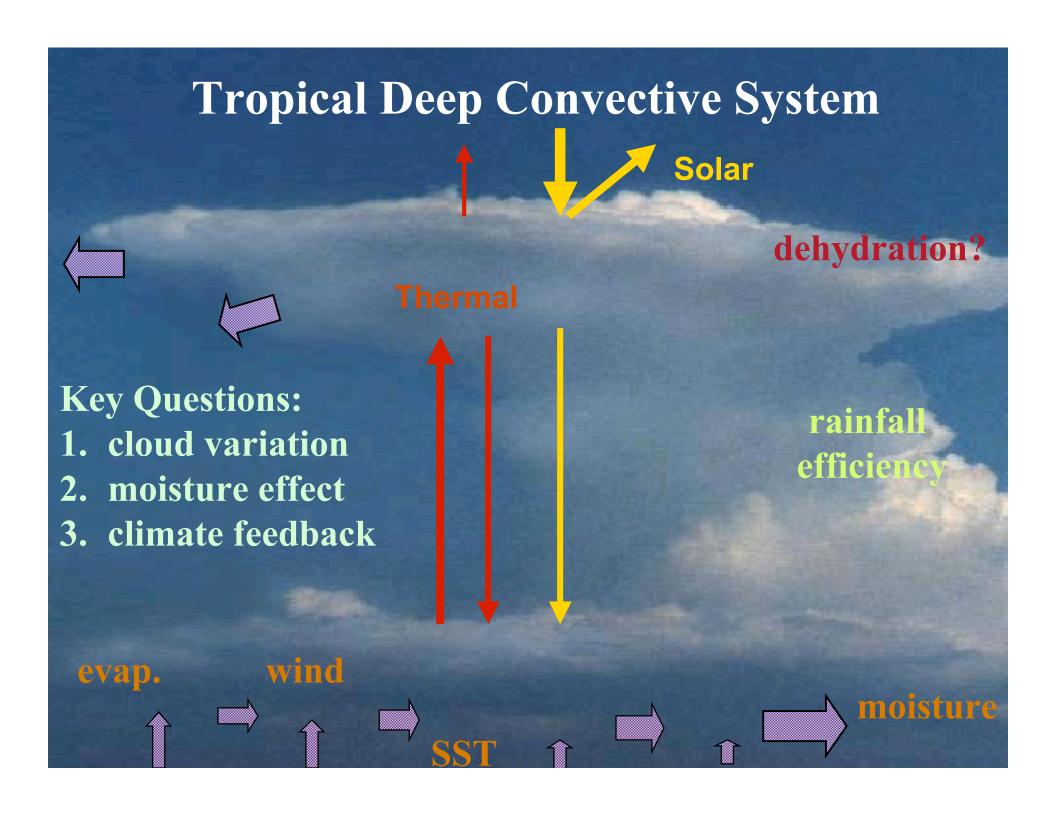


# Climate feedback of tropical deep convection — moisture budget analysis

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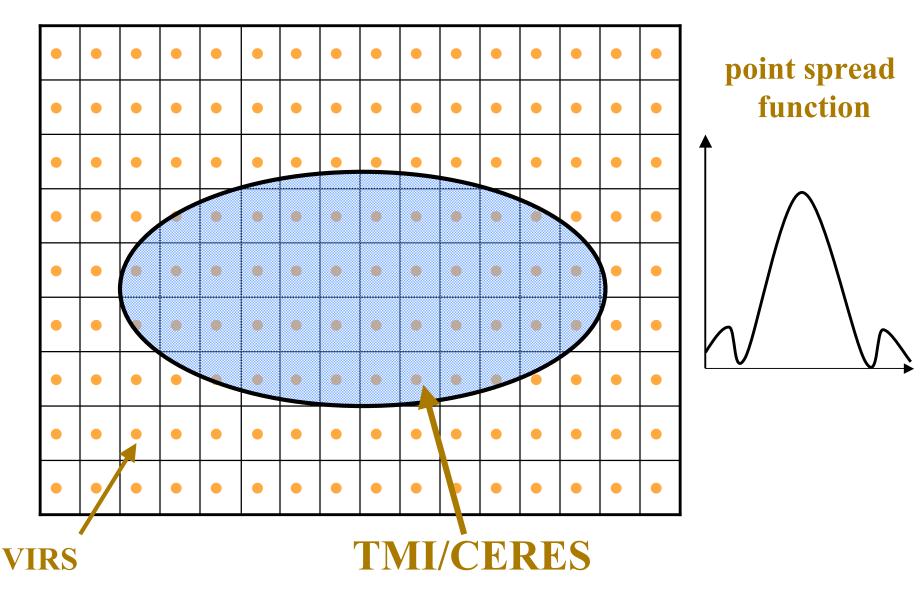
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# Data match —— convolution







#### Multi-Sensor Data Fusion



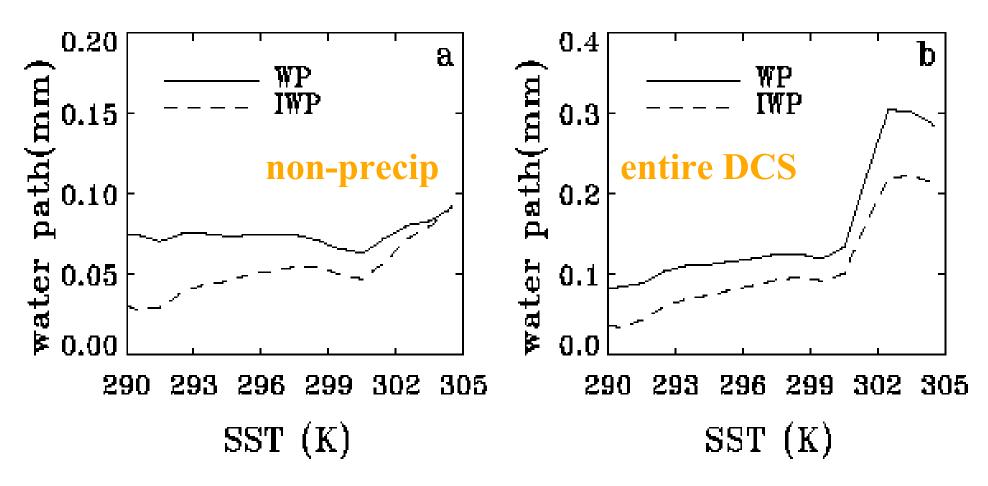
- Convolution of multi-sensor measurements into TMI 37GHz FOV
- · Rainfall rate: TMI & PR
- Cloud detection, top temperature and optical depth: VIRS
- · Cloud liquid/ice water path: VIRS and TMI
- Evaporation and boundary layer moisture: TMI
- Radiation fields: CERES

· Vector wind fields: assimilation products



#### DCS WP & IWP

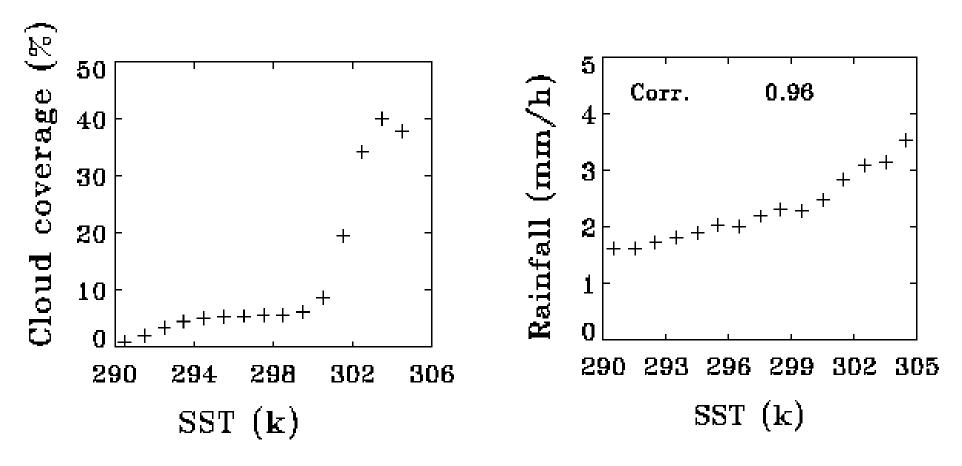




Deeper convection under higher SST conditions generally produces thicker anvils, and is likely moistening upper troposphere more.



# DCS areal coverage & rain rate

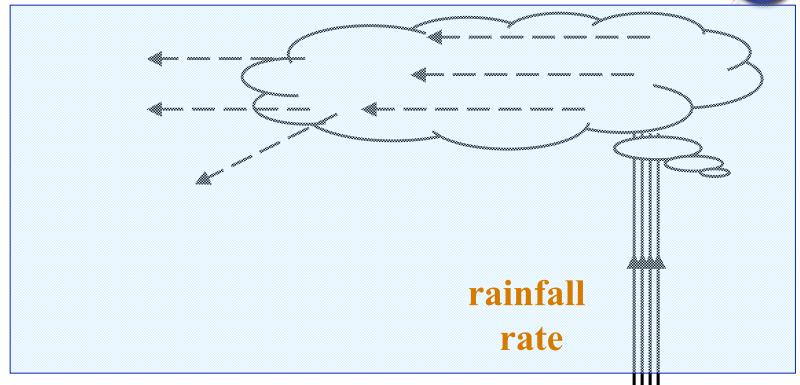


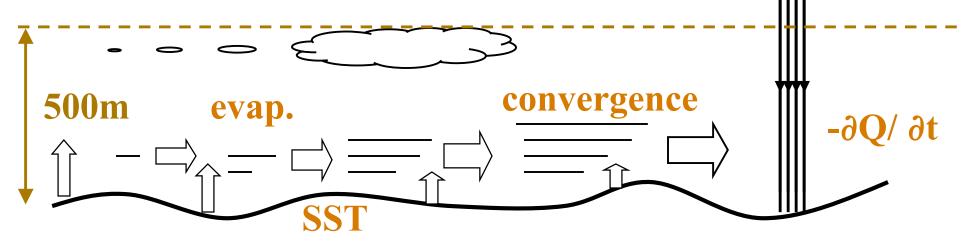
Deeper convection under higher SST conditions could cause higher rainfall efficiency, which would result in DCS reductions in warmer environments. Why is there no evidence of DCS dehydrations?



### Moisture Balance



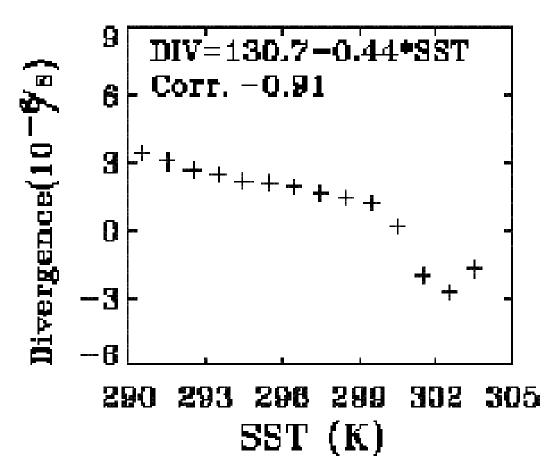






# divergence





**Convergence:** sharp increases in warm SSTs



#### Rainfall efficiency



Moisture supply to DCS clusters:

$$\mathbf{M}_{s} = -\partial \mathbf{Q}_{b} / \partial t + \mathbf{E}_{s} - \nabla \cdot \mathbf{Q}_{b} \mathbf{V}$$
 tendency, evaporation and horizontal transport

•  $\eta = RR/Ms$ 

$$= RR/(-\partial WV_L/\partial t + E_s - \nabla \cdot WV_LV)$$
 (2)

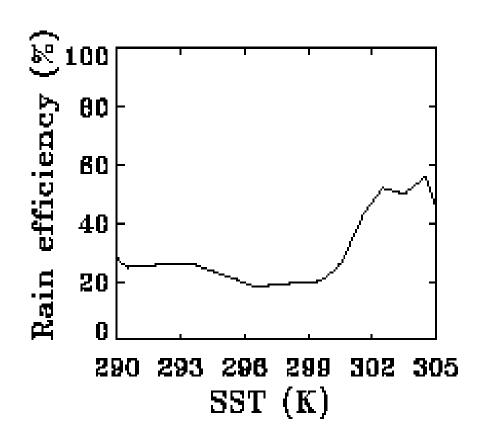
TMI: Es and WV<sub>L</sub> for lowest 500m PBL

ECMWF:  $\nabla \cdot WV_LV$ 





## Rainfall efficiency



mean: ~35%

increase: ~ 2%/K

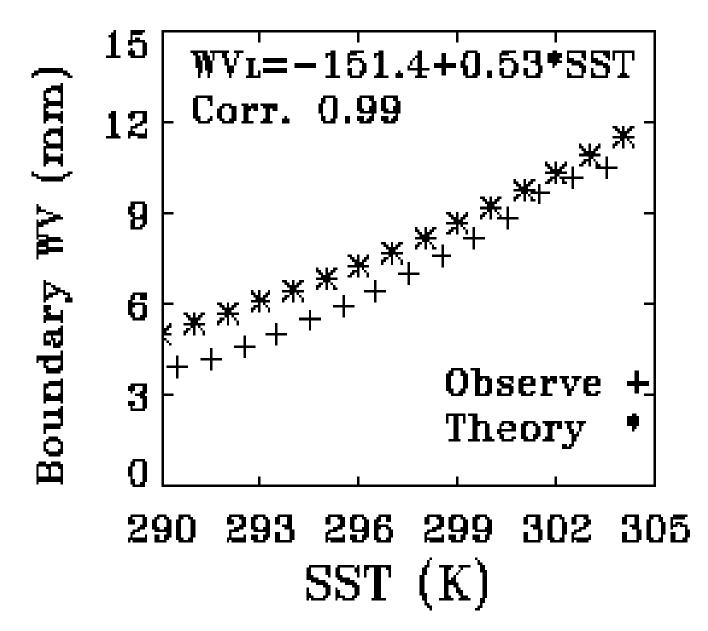
sharp for warm SST

Why is there no evidence of DCS dehydrations?



## Boundary layer moisture







## Moisture change with Temp



• 
$$Q_b$$
 = rh \*  $Q_{sa}$   
 $Q_{sa}$ : saturated Q; rh: relative humidity  
 $\Delta Q_b$  = rh \* 6.3%  $Q_{sa}$  = 6.3%  $Q_b$  (3)

• 
$$\mathbf{E}_{s} = \rho_{a} \mathbf{C}_{l} (\mathbf{Q}_{ss} - \mathbf{rh} * \mathbf{Q}_{sa}) \mathbf{W}_{s}$$
  

$$\Delta \mathbf{E}_{s} = 6.3\% \mathbf{E}_{s}$$
(4)

So, all moisture related terms increase ~6.3%/K.



## Moisture transport for anvils



•  $M_{cld} = Ms - RR = Ms (1 - \eta)$  (6) where  $M_{cld}$  is moisture supply to cirrus-anvil clouds

• 
$$\Delta M_{cld} = \Delta (Ms (1 - \eta))$$
 (7)

•  $\Delta M_{eld}/M_{eld} = \Delta M_{eld}/(Ms(1-\eta))$ 

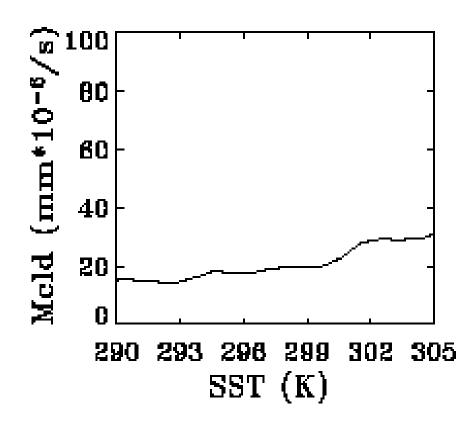
$$= \Delta Ms/Ms - \Delta \eta/(1 - \eta)$$

•  $\approx 6.3\% - 2\%/(1 - 0.35) \approx 3.0\%$  (8)





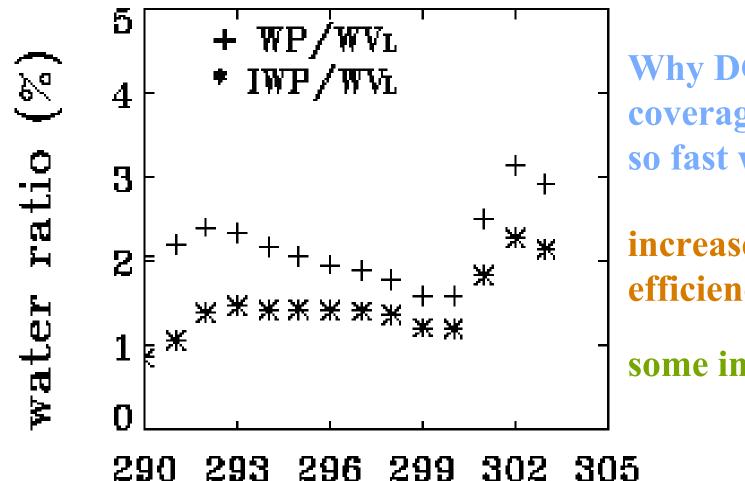






#### Water ratio





Why DCS area coverage increase so fast with SST?

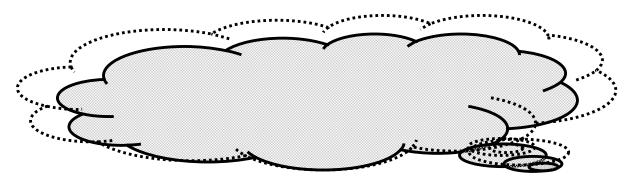
increase in cloud efficiency?

some indication



# Generalized radiative forcing





clear areas

existing environmental conditions no high clouds



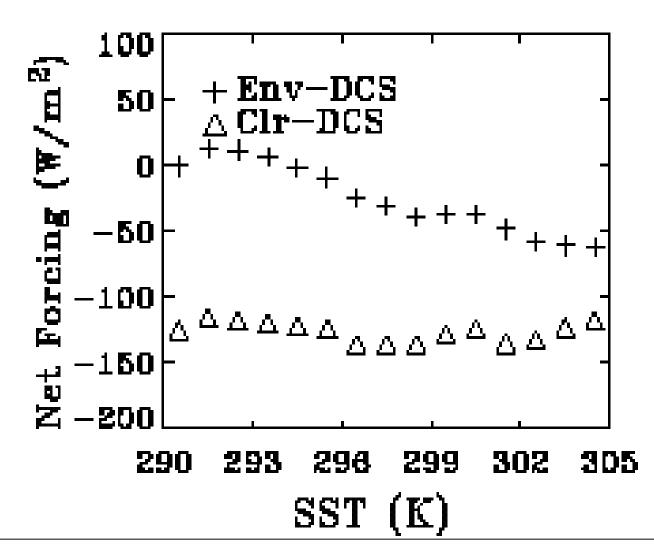
low clouds

SST



#### Net radiation





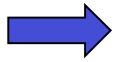
Note: Clr, Env and DCS represent values for clear, environmental and DCS skies. Thus, F = Clr - DCS; G = Env - DCS



#### **Summary**



- Although DCS precipitation and rainfall efficiency increase with temperature, DCS area coverage still increases with SST.
- The boundary layer moisture supply for DCS increases faster than rainfall efficiency and results in increases of the moisture transported to the upper troposphere for cirrus-anvil formation.
- The average change in net radiative forcing due to DCS change from existing environmental conditions is relatively weak (about –0.56 W/m²/K).



a weak negative DCS feedback





#### Acknowledgement

Many people, especially Sunny Sun-Mack, Jianping Huang, Dave Young, and Gary Gibson have significant contributions to this study.

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